Environmental Geochemistry

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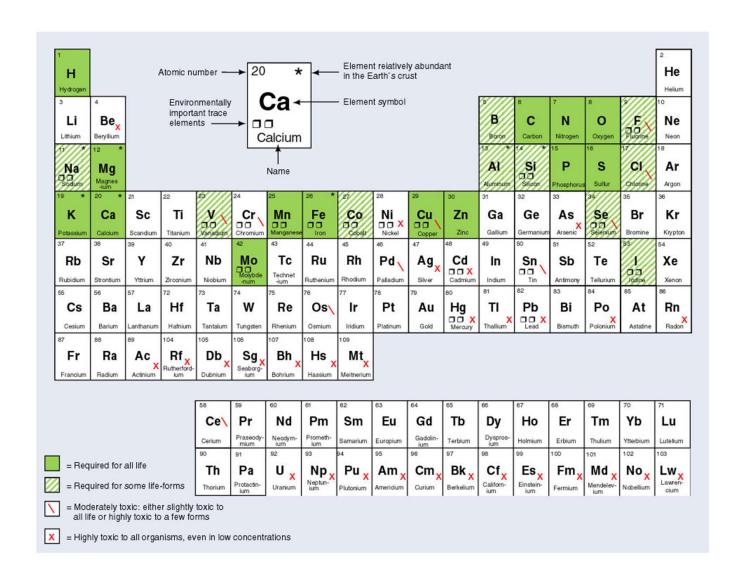
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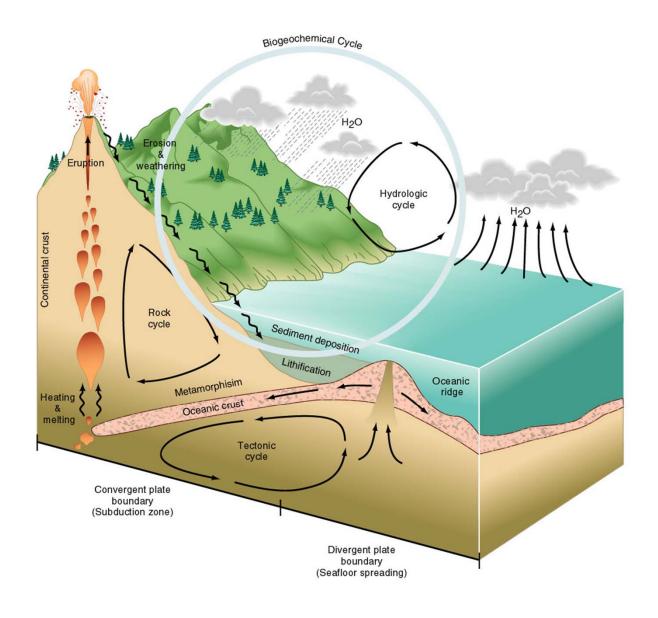
- Time: 8:00-9:00
- · Day: Sunday, Tuesday, Thursday
- · Room: GEO106
- · Hours Credit: 3
- Textbook: Nelson Eby (2004):
 Principles of Environmental Geochemistry

Environmental geochemistry

 Is the application of chemical principles to predicting the fate of organic and inorganic pollutants at the earth's surface and in the atmosphere

Fig 4.5 Periodic table of the elements.





Idealized diagram of the geologic cycle, which includes the tectonic, hydrologic, rock and biogeochemical cycles.

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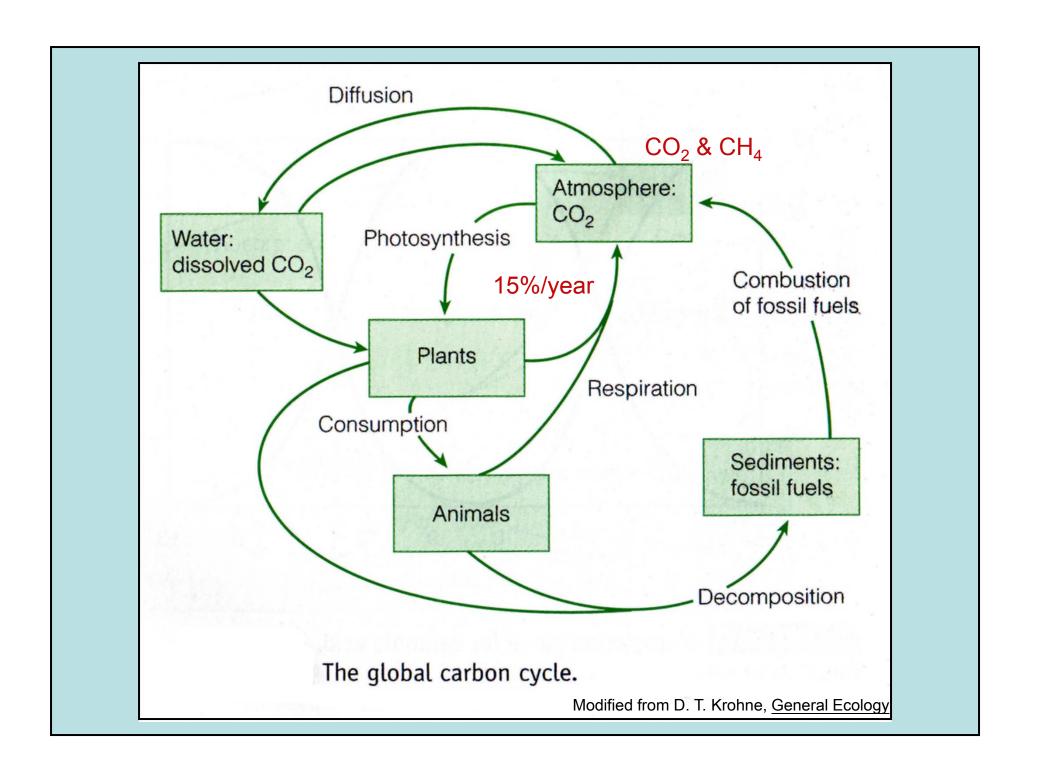


Fig 4.14 The Carbon Cycle.

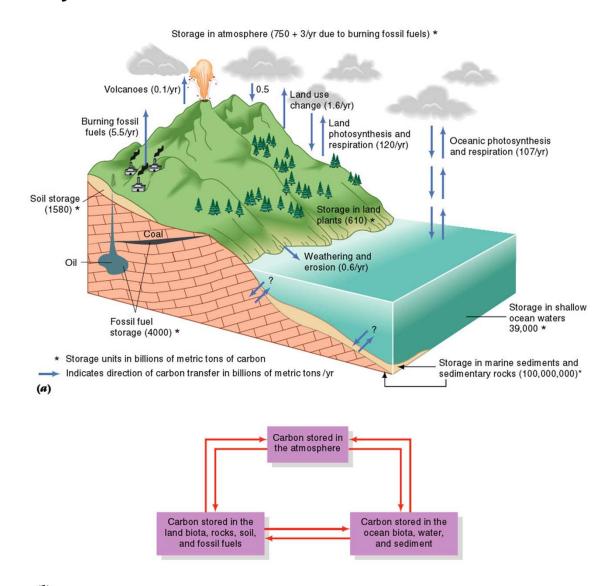
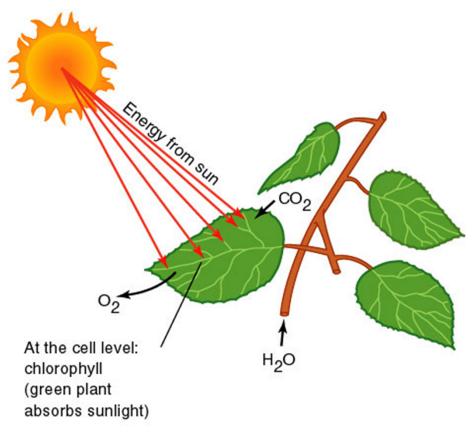


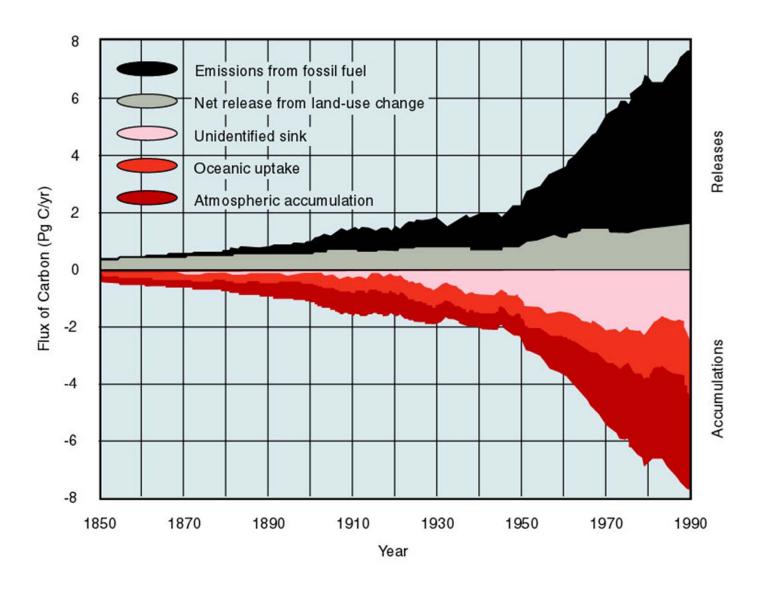
Fig 4.15 Idealized diagram illustrating photosynthesis for a green plant (tree) and generalized reaction.



General Photosynthesis: chemical reaction

$$6CO_2 + 6H_2O \xrightarrow{\text{sunlight}} C_6H_{12}O_6 + 6O_2$$
carbon + water $\xrightarrow{\text{sunlight}}$ sugar (glucose) + oxygen

Fig 4.17 Global flux of carbon, 1850-1990.



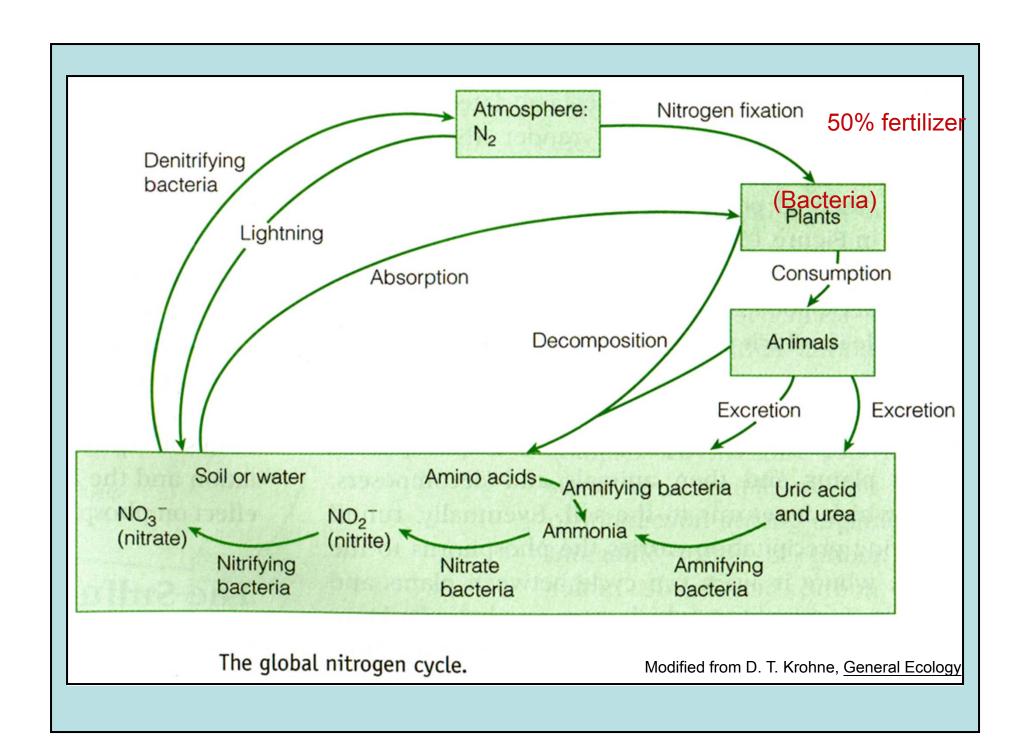
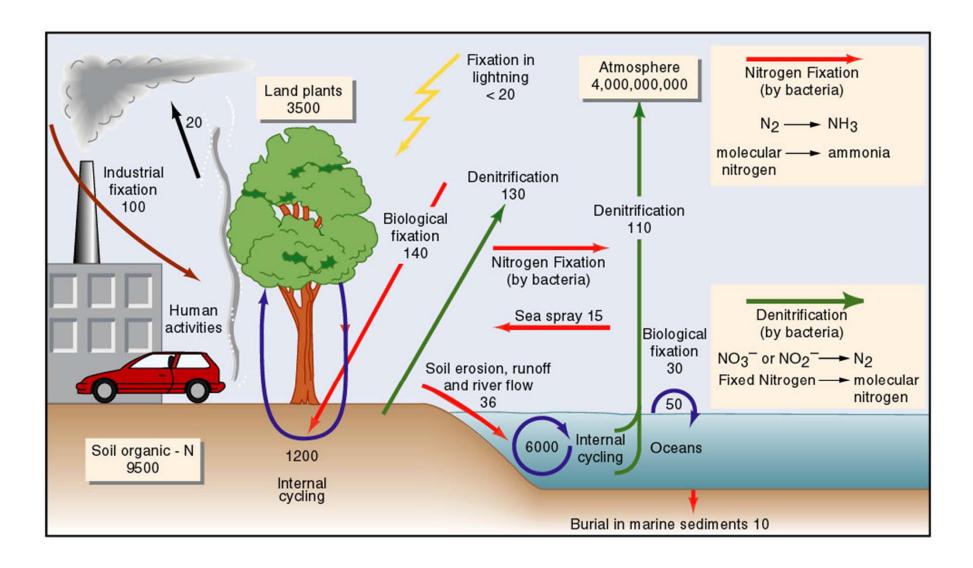
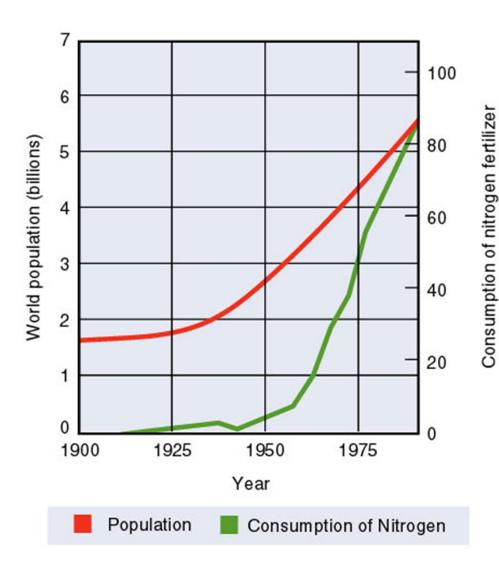


Fig 4.19 The global nitrogen cycle.



Critical Thinking Issue – How are Human Activities Affecting the Nitrogen Cycle?



(megatons of nitrogen) Effects of increased use of

nitrogen fertilizer:

- Increased nitric acid in soil:
 - Leaching of magnesium and potassium
 - Increased aluminum levels
 - Plant root damage
 - Changes in microbe communities
 - Fish kills
- Eutrophication of water bodies
- Nitrates in drinking water
- Global effects on plant based CO₂ uptake

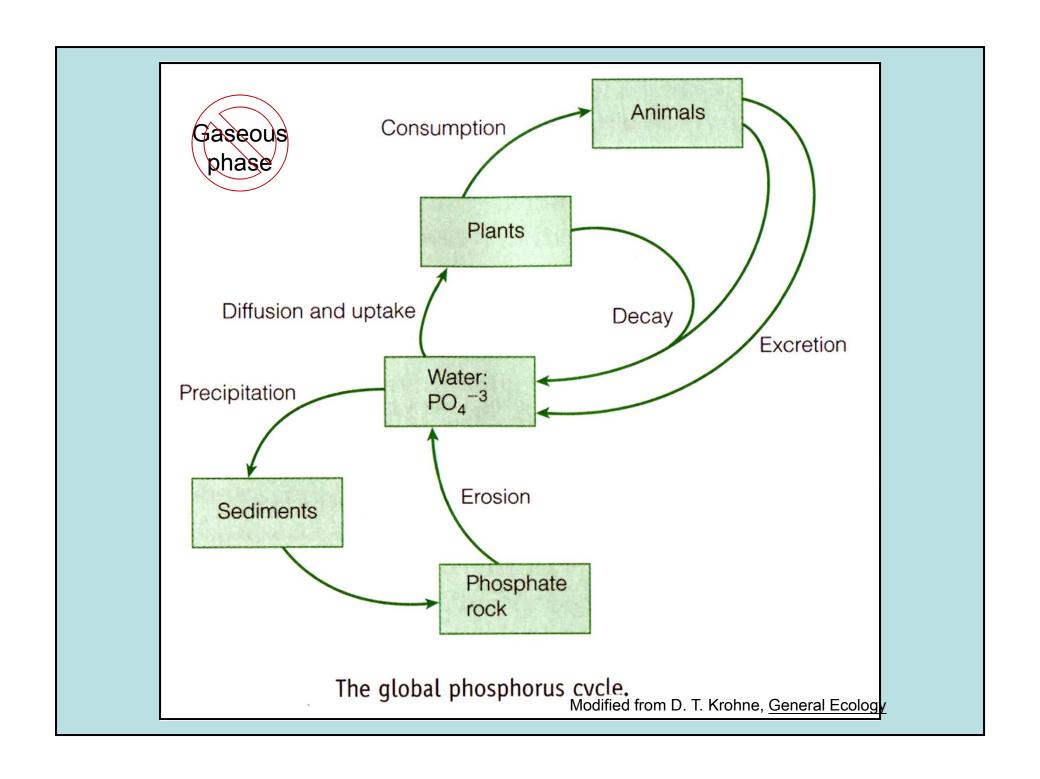
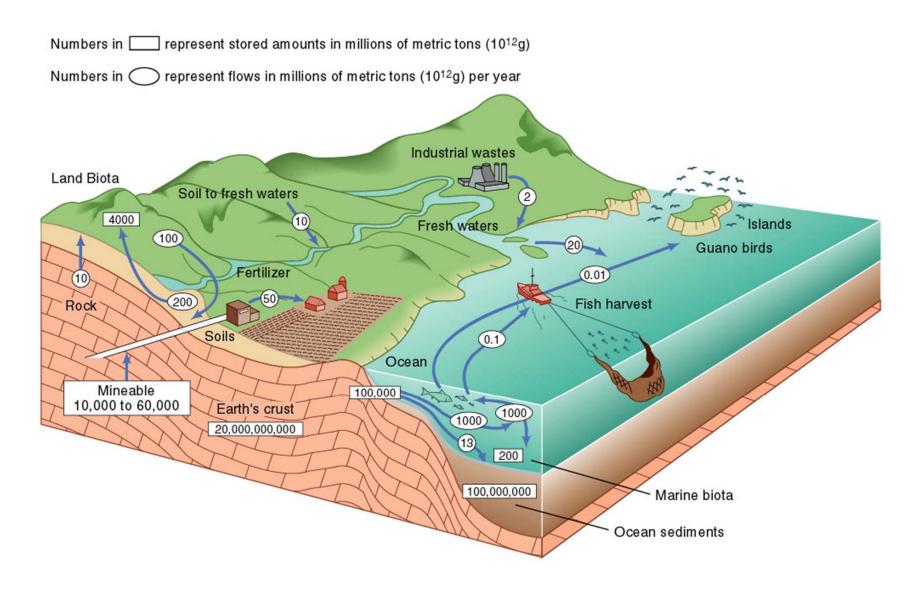
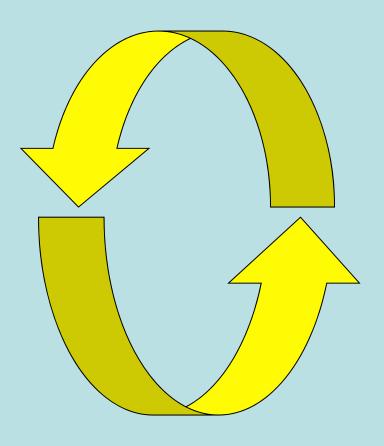


Fig 4.20 The global phosphorus cycle.



Ecosystems and the Gaia Hypothesis



Ecosystem defined: a community of organisms and it's corresponding abiotic environment through which matter cycles and energy flows

- Wide variation in ecosystems
- Boarders can be well defined or vague
- Can be natural or artificial, managed or wild
- Wide range in scale
- Common to all ecosystems: energy flow and cycling of matter

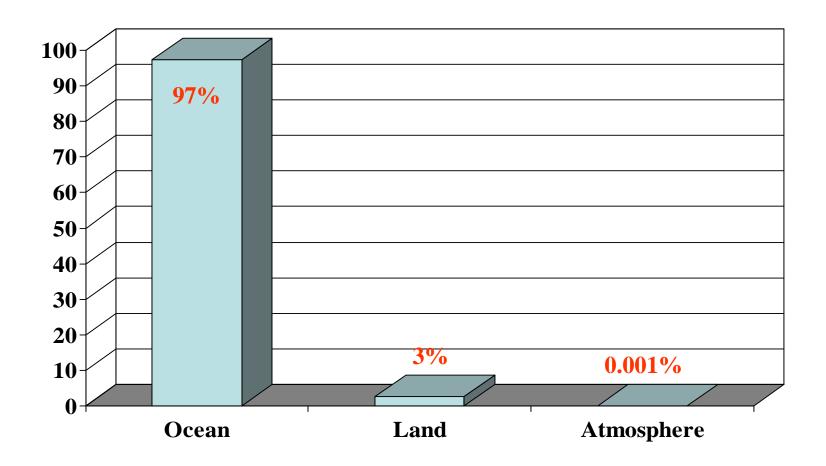
Lecture Overview

- Properties of Water
 - Physical
 - Chemical

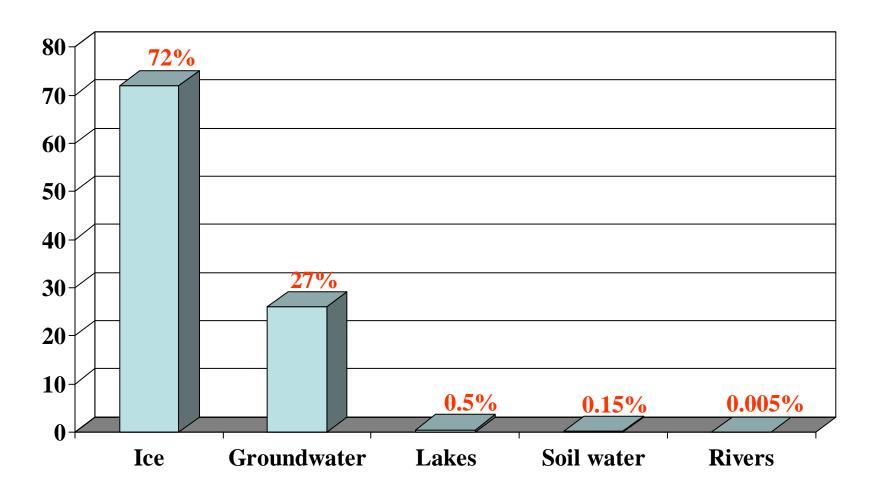
Did you know...

- ☐ The earth is the only planet in the solar system where water occurs in all three phases (solid, liquid, gas)
- □ Water covers 70% of the earth's surface
- ☐ The amount of water (1.4 billion km³) in the earth system is constant
- ☐ The oceans accounts for over 97% of the earth's water
- □ Water is in constant movement in a circulation called the hydrologic cycle
- □ 60-75% of the human body is water

Occurrence of water on the earth



Occurrence of non-oceanic water



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Classification of Waters

- Meteoric recently derived from the atmosphere as rain or snow
- Formation pore water in deep sedimentary rocks
- Metamorphic from dehydration reactions
- · Juvenile from the mantle

Table 1-7. Properties of water

Property	Comparison to other substances
Heat capacity	Highest of all common liquids (except ammonia) and solids
Latent heat of fusion	Highest of all common liquids (except ammonia) and most solids
Latent heat of vaporization	Highest of all common substances
Dissolving ability	Dissolves more substances (particularly ionic compounds), and in greater quantity than any other common liquid.
Transparency	Relatively high for visible light
Physical state	Water is the only substance that occurs naturally in all three states at the earth's surface
Surface tension	Highest of all common liquids
Conduction of heat	Highest of all common liquids (Hg is higher)
Viscosity	Relatively low viscosity for a liquid

METALS VS. NON-METALS

- Ionization potential (IP) The energy required to remove an electron from a neutral atom to an infinite distance from the nucleus.
- Electron affinity (EA) The tendency of a neutral atom to attract electrons.
- Metal An element with a low IP and a low EA; it tends to lose electrons and form a cation.
- Nonmetal An element with a high IP and high EA; it tends to gain electrons and form an anion.

IONIC BONDS

- <u>Ionic bonds</u> bonds formed via the electrostatic attraction between oppositely charged ions.
- Coulomb's law $F = \frac{Z_{+}Z_{-}}{4\pi\epsilon_{0}r^{2}}$

 ε_0 = permittivity of free space = 8.84x10⁻⁹ farads m⁻¹

1 farad = 1 coulomb volt⁻¹

 $F = force of attraction; Z_+ and Z_- are the charges of the cation and anion respectively; r is the cationanion distance.$

COVALENT BONDS

- <u>Covalent bonds</u> bonds formed by sharing of electrons.
- Many gaseous species and organic compounds contain covalent bonds, e.g., N₂.
- A few minerals, such as diamond and graphite, contain covalent bonds.
- Sharing may not be equal. That is, covalent bonds may have some ionic character, and vice versa.

MEASURE OF IONIC/COVALENT CHARACTER

- Electronegativity (χ) measures the ability of an atom in a compound to attract electrons to itself.
- Metals have low χ values and non-metals have high χ values (see Table 1-2 in Kehew for electronegativity values of the elements).
- We use differences in electronegativity to determine degree of ionic character of a bond.

■ Example: CsF
$$\chi_{Cs} = 0.7; \ \chi_{F} = 4.0$$
 $\chi_{F} - \chi_{Cs} = 3.3$

> 92% ionic; < 8% covalent

TABLE 1-2 Electronegativity values of the elements.

H 2.1 Li																	He
Li 1.0	Be 1.5											B 2.0	C 2.5	N 3.0	O 3.5	F 4.0	Ne
Na 0.9	Mg 1.2											Al 1.5	Si 1.8	P 2.1	S 2.5	Cl 3.0	Ar
K 0.8	Ca 1.0	Sc 1.3	Ti 1.5	V 1.6	Cr 1.6	Mn 1.5	Fe 1.8	Co 1.8	Ni 1.8	Cu 1.9	Zn 1.6	Ga 1.6	Ge 1.8	As 2.0	Se 2.4	Br 2.8	Kr
Rb 0.8	Sr 1.0	Y 1.2	Zr 1.4	Nb 1.6	Mo 1.8	Tc 1.9	Ru 2.2	Rh 2.2	Pd 2.2	Ag 1.9	Cd 1.7	In 1.7	Sn 1.8	Sb 1.9	Te 2.1	I 2.5	Xe
Cs 0.7	Ba 0.9	La-Lu 1.1-1.2	Hf 1.3	Ta 1.5	W 1.7	Re 1.9	Os 2.2	Ir 2.2	Pt 2.2	Au 2.4	Hg 1.9	T1 1.8	Pb 1.8	Bi 1.9	Po 2.0	At 2.2	Rn
Fr 0.7	Ra 0.9	Ac-Lr 1.1															

PERCENT IONIC CHARACTER

Electronegativity difference 0.1	% Ionic character 0.5	Electronegativity difference 1.2	% Ionic character 30	Electronegativity difference 2.3	% Ionic character
0.2	1	1.3	34	2.4	76
0.3	2	1.4	39	2.5	79
0.4	4	1.5	43	2.6	82
0.5	6	1.6	47	2.7	84
0.6	9	1.7	51	2.8	86
0.7	12	1.8	55	2.9	88
0.8	15	1.9	59	3.0	89
0.9	19	2.0	63	3.1	91
1.0	22	2.1	67	3.2	92
1.1	26	2.2	70	>3.2	>92

WHY DO WE CARE IF A BOND IS IONIC OR COVALENT?

- All physical and chemical properties of a compound depend on the character of the bonds.
- Example: Solubility in water

A general rule is that, like dissolves like. As we will see, water is a polar covalent solvent.

Ionically bonded compounds have high aqueous solubilities

e.g., NaCl
$$\chi_{Cl}$$
 - χ_{Na} = 3.0 - 0.9 = 2.1

67% ionic, relatively soluble in water

C(diamond)
$$\chi_C - \chi_C = 0$$

0% ionic, nearly insoluble in water

RULES TO DETERMINE OXIDATION STATES

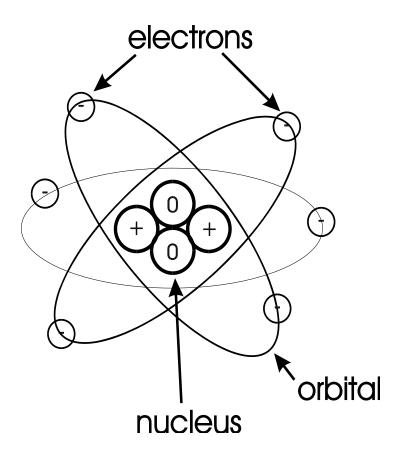
- 1) The oxidation state of all pure elements is zero.
- 2) The oxidation state of H is +1, except in hydrides (e.g., LiH, PdH₂), where it is -1.
- 3) The oxidation state of O is -2, except in peroxides (e.g., H_2O_2), where it is -1.
- 4) The algebraic sum of oxidation state must equal zero for a neutral molecule or the charge on a complex ion.

ELEMENTS WITH VARIABLE OXIDATION STATES

- Sulfur: $SO_4^{2-}(+6)$, $SO_3^{2-}(+4)$, S(0), $FeS_2(-1)$, $H_2S(-2)$
- Carbon: $CO_2(+4)$, C(0), $CH_4(-4)$
- Nitrogen: $NO_3^-(+5)$, $NO_2^-(+3)$, NO(+2), $N_2O(+1)$, $N_2(0)$, $NH_3(-3)$
- Iron: $Fe_2O_3(+3)$, FeO(+2), Fe(0)
- Manganese: $MnO_4^-(+7)$, $MnO_2(+4)$, $Mn_2O_3(+3)$, MnO(+2), Mn(0)
- Copper: CuO(+2), $Cu_2O(+1)$, Cu(0)
- Tin: $SnO_2(+4)$, $Sn^{2+}(+2)$, Sn(0)
- Uranium: $UO_2^{2+}(+6)$, $UO_2(+4)$, U(0)
- Arsenic: $H_3AsO_4^0(+5)$, $H_3AsO_3^0(+3)$, As(0), $AsH_3(-1)$
- Chromium: $CrO_4^{2-}(+6)$, $Cr_2O_3(+3)$, Cr(0)
- Gold: $AuCl_4^-(+3)$, $Au(CN)_2^-(+1)$, Au(0)

ATOMIC STRUCTURE

- Each atom of an element is composed of a nucleus surrounded by electrons in various orbitals.
- Nucleus a central concentration of mass consisting of protons and neutrons.
- Electrons negatively charged particles of relatively low mass.
- Protons positively charged particles of relatively high mass.
- Neutrons particles with no charge but mass similar to that of protons.



CHEMICAL BONDING

 lonic bonds - bonds formed via the electrostatic attraction between oppositely charged ions.

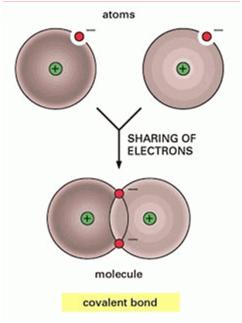
$$Na \cdot + \cdot CI \longrightarrow Na^{\dagger} + \cdot CI$$

Na contributes electron, leaving it with a closed shell Cl gains electron, leaving it with a closed shell

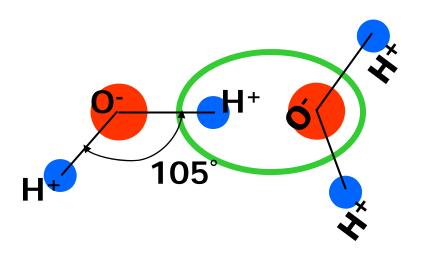
Covalent bonds - bonds formed by sharing of electrons.

$$H \bullet + \bullet H \longrightarrow H^*_{\bullet}H$$

Polar Covalent bonds - the sharing of the electron pair is unequal. There is a charge separation with one atom being slightly more positive and the other more negative.

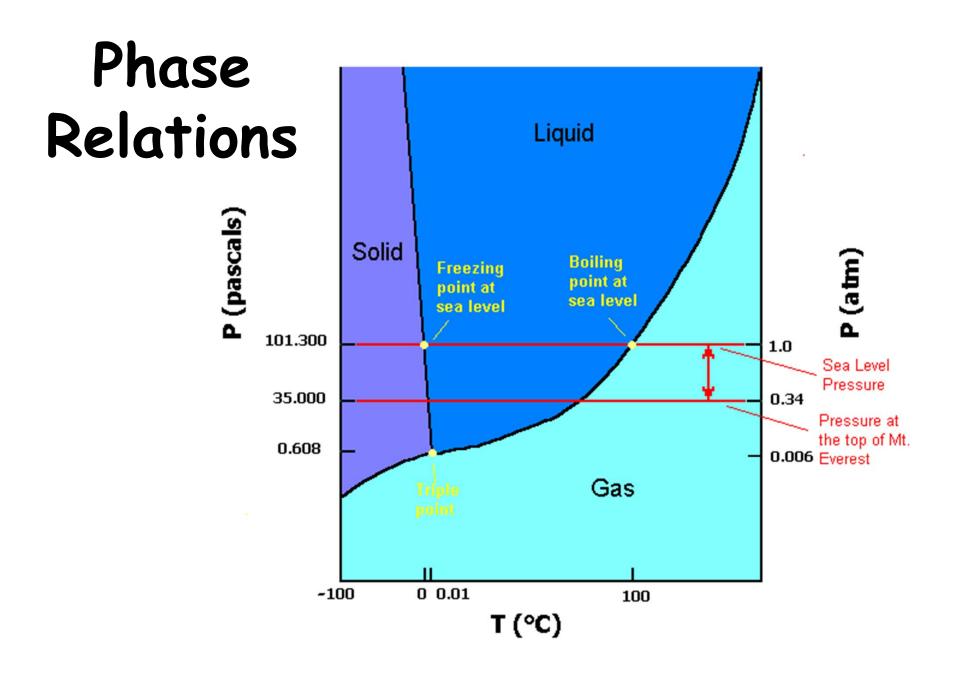


PROPERTIES OF WATER

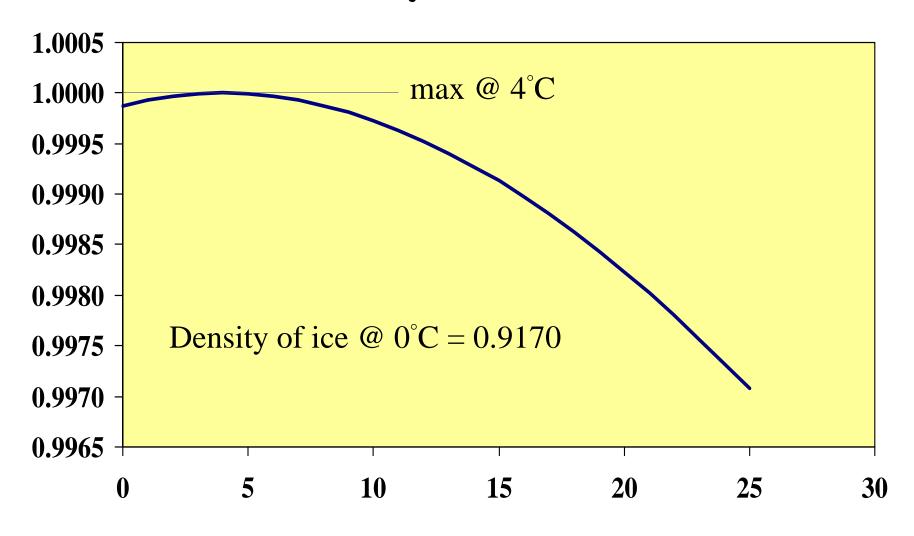


Because each water molecule has a positive and a negative end, these can attract one another to form a *hydrogen* bond.

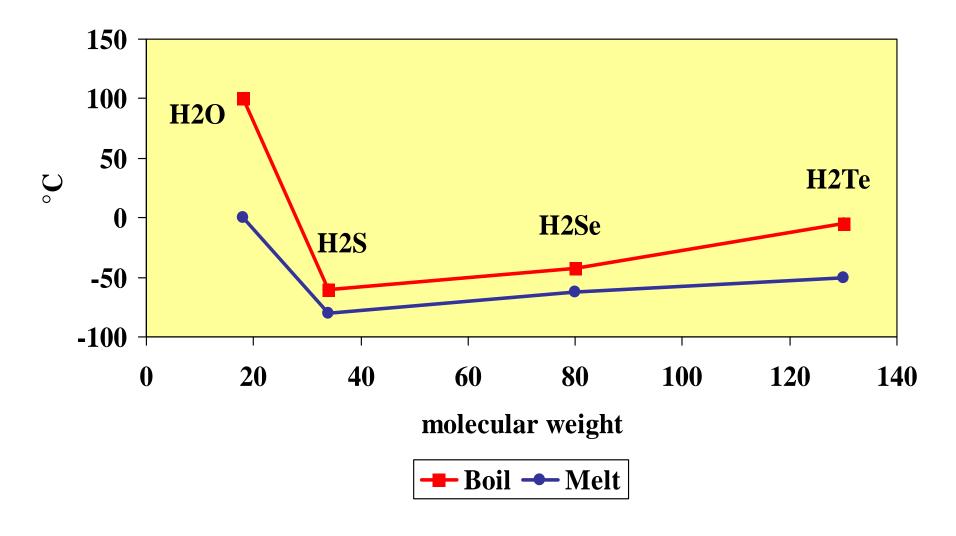
- The hydrogen bonding of water is responsible for many anomalous properties of water.
 - high boiling point & melting point
 - The maximum density is at 4C.
 - Ice is less dense than the liquid water.



Density of Water

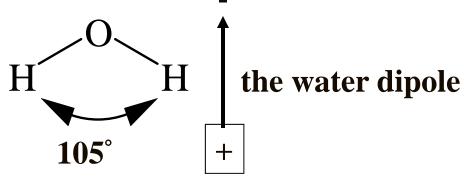


Boiling & Melting Points



WHY DO WE CARE IF A BOND IS IONIC OR COVALENT?

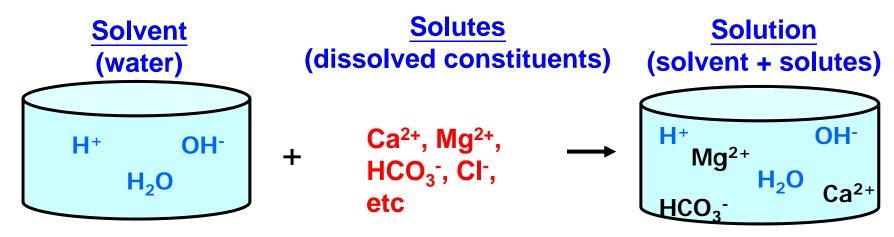
- All physical and chemical properties of a compound depend on the character of the bonds.
- Solubility in water (general rule like dissolves like)
 Water is a polar covalent solvent 39% ionic & 61% covalent meaning there is a positive and a negative pole to the molecule.



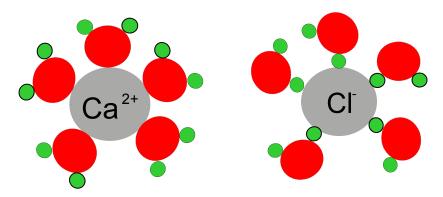
Water dissolves ionic solids (e.g., NaCl, K_2SO_4) very well. On the other hand, polar water is a very poor solvent for many organic compounds, e.g., toxic benzene (C_6H_6).

PROPERTIES OF WATER

Solvent: Water dissolves some amount of virtually every solid or gas.



- Also because of the polar nature of water, ions will be surrounded by water dipoles (hydrated) in solution.
- Hydration isolates the ions from their neighbors and neutralizes the attractive forces that hold minerals together.



UNITS OF EXPRESSION

Mass Concentration - we use many units of concentration.

weight/weight units – ppm, ppb

$$ppm = \frac{\text{mass of solute (mg)}}{\text{mass of solution (kg)}}$$

• weight/volume units – mg/L, g/L $mg/l = \frac{mass \text{ of solute } (mg)}{volume \text{ of solution } (l)}$

The conversion factor between mg/L and ppm is:

$$\frac{ppm}{l} (parts per million) = \frac{1g}{10^6 g} = \frac{1mg}{lkg} = \frac{1mg}{l} (in natural water)$$

$$\frac{mg}{kg} = \frac{mg}{l} \times \frac{1}{\rho}$$

Because the density of many natural waters is near 1 kg/L, it is often a sufficiently good approximation that mg/L and ppm are numerically equal.

ppm (parts per million) =
$$\frac{1g}{10^6 g} \times \frac{10^3 mg}{1g} \times \frac{10^3 g}{1 kg} = \frac{1 mg}{1 kg}$$

Molar Concentration - In almost all geochemical calculations, it is necessary to use molar concentrations rather than mass concentrations.

Moles (1 mole = 6.023×10^{23} atoms or molecules)

• Molality (m) =
$$\frac{\text{moles of solutes, mol}}{\text{mass of solvent, kg}}$$

• Molarity (M) =
$$\frac{\text{moles of solutes, mol}}{\text{volume of solution, L}}$$

- Molarity is the most common concentration unit involved in calculations dealing with volumetric stoichiometry.
- Conversion from mol/L (M) to mg/L is accomplished using the formula:

$$\frac{mg}{L} = \frac{mol}{L} \times \underline{molar\ mass} \left(\frac{g}{mol} \right) \times \frac{1000\ mg}{g}$$

Example. What's the molar mass for 1 mole of HCI?

1 1 H Hydrogen 1.00794	2 IIA	13 IIIA	14 IVA	15 VA	16 VIA	17 VIIA	2
3 2 Li Lithium 6.941	4 Be Beryllium 9.012182	5 2 B Boron 10.811	6 2 4 C Carbon 12.0107	7 2 5 N Nitrogen 14.00674	O Owgen Average	9 ² 7 F Eluorine Je Atomic	Ne Neon
11 2 8 Na 1 Sodium 22.989770	12 2 8 2 Mg 2 Magnesium 24.3050	13 2 Al 3 Aluminum 26.981538	14 2 8 3 Si 4 Silicon 28.0855	15 2 8 5 Phosphorus 30.973761	16 S Sulfur 32.066	17 2 8 7 Cl 7 Clylorine 14527	18 2 Ar 8 Argon 39.948
19 2 8 8 8 Potassium 39.0983	20 2 8 8 Calcium 40.078	31 2 8 Ga 18 Gallium 69.723	32 2 8 Ge 18 Germanium 4 72.61	33 2 8 As 18 Arsenic 5 74.92160	34 2 8 18 18 Selenium 78.96	35 2 Br 18 Bromine 79,904	36 2 Kr 18 Krypton 83.80
37 2 8 8 18 Rubidium 1 85.4678	38 2 8 Sr 18 Strontium 8 87.62	49 2 8 18 18 Indium 3 114.818	50 2 8 Sn 18 Tin 18 118,710	51 2 8 Sb 18 Antimony 5 121.760	52 2 8 18 18 Tellurium 6 127.60	53 2 8 18 18 18 126.90447	54 2 Xe 18 Xenon 18 131.29
55 2 8 Cs 18 Cesium 8 132,90545 1	56 2 8 8 18 18 18 18 137.327 2	81 2 8 TI 18 Thallium 32 18 204.3833 3	82 2 Pb 18 Lead 32 18 207.2 4	83 2 Bi 18 Bismuth 32 18 208.98038 5	84 2 PO 18 Polonium 32 Polonium 18 (209) 6	85 2 At 18 Astatine 32 Astatine 18 (210) 7	86 2 Rn 18 Radon 18 (222) 8

H: 1g (atomic mass) x 1 = 1g
$$1 + 35.5 = 36.5$$
 g/mole CI: 35.5g (atomic mass) x 1 = 35.5g

Mole fraction (X) - another form of molar concentration

$$X_{Solute} = \frac{moles of solute}{total moles of all components}$$

the mole fraction (X) is used for solid solutions, e.g., solid solutions between $KAISi_3O_8$ and $NaAISi_3O_8$. In such a solid solution, the mole fraction of $KAISi_3O_8$ would be written as:

$$\mathbf{X}_{\mathsf{KAISi_3O_8}} = \frac{\mathsf{moles}\,\mathsf{KAISi_3O_8}}{\mathsf{moles}\,\mathsf{KAISi_3O_8} + \mathsf{moles}\,\mathsf{NaAISi_3O_8}}$$

• Percentage (%) - the ratio of a solute to the solution If a solution concentration is given as a percentage, you can generally assume it is a mass percentage unless otherwise stated.

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Example. Suppose you are given a concentrated solution of HCl which is known to be 37.0% HCl and has a solution density of 1.19 g/mL. What is the Molarity of HCl?

(1) Begin with the assumption of 100 g of solution.

(2) To find molarity, we need to determine the moles of HCI (solute).

moles (mol) =
$$\frac{\text{mass (g)}}{\text{molar mass (g/mol)}} = \frac{37g}{36.5g/\text{mol}} = 1.01 \text{mol}$$

(3) Convert the known mass of solution, 100 g solution, to liters of solution, using the density of the solution:

$$\frac{100 \text{ g}}{1.19 \text{ g/mL}} \times \frac{1 \text{L}}{1000 \text{ mL}} = 0.084 \text{ L solution}$$

(4) Calculate the molarity (M)

$$\frac{1.01\,\text{mol HCI}}{0.084\,\text{L solution}} = 12.0\,\text{mol/L}$$

UNITS OF EXPRESSION

Equivalents and Normality (N)- units: equivalents/liter

Equivalents (eq) are similar to moles, but take into account the valence of an ion.

0.002 mol L⁻¹ of Ca²⁺ = 0.004 eq L⁻¹ Ca²⁺ 0.001 mol L⁻¹ of Na⁺ = 0.001 eq L⁻¹ Na⁺ 0.003 mol L⁻¹ Al³⁺ = 0.009 eq L⁻¹ Al³⁺

- ♣ Normality (N) is another name for eq L⁻¹.
- ♣ Alkalinity is an important solution parameter that is expressed as eq L⁻¹ or meq L⁻¹. Hardness is another parameter expressed as eq L⁻¹.

Example. The laboratory reported the concentration of Ca²⁺ in a water sample as 92 mg/L. What is the normality of Ca²⁺? **Valence (charge)**

$$\frac{\text{Equivalent}}{\text{Liter}} = \text{Conc.}(\frac{\text{mg}}{\text{L}}) \times \frac{1 \text{mole}}{\text{atomic or molecular wt.}} \times \frac{\text{\# of equivalents}}{\text{mole}}$$

$$92\frac{mg}{L} \times \frac{1 mole}{40.08g} \times \frac{1g}{1000mg} \times \frac{2equiv}{mole} = 4.6 \times 10^{-3} \frac{equiv}{L} = 4.6 \frac{milliequiv}{L}$$

Box Models and Geochemical Cycles

- Box models and the steady-state assumption are widely used in the environmental sciences to trace the passage of various constituents through the bio-geo-hydro-atom-spheres and to assess the impact of anthropogenic inputs on these natural cycles.
- A box model consists of several boxes showing the reservoirs for a particular substance and the rate at which material is transferred between the reservoirs.

- In steady-state system, the total amount of a substance in each reservoir remains constant, thus the rate of addition to, and removal of a material from a reservoir must be the same.
- The hydrologic cycle box is the simplest box model shown in the next slide.
- From the figure, we can calculate the residence time for water in the various reservoirs.
- Residence time: is the average length of time a particular substance will reside in a reservoir.

Residence Time =
$$\frac{\text{Amt. of material in reservior}}{\text{Rate of addition (removal)}}$$

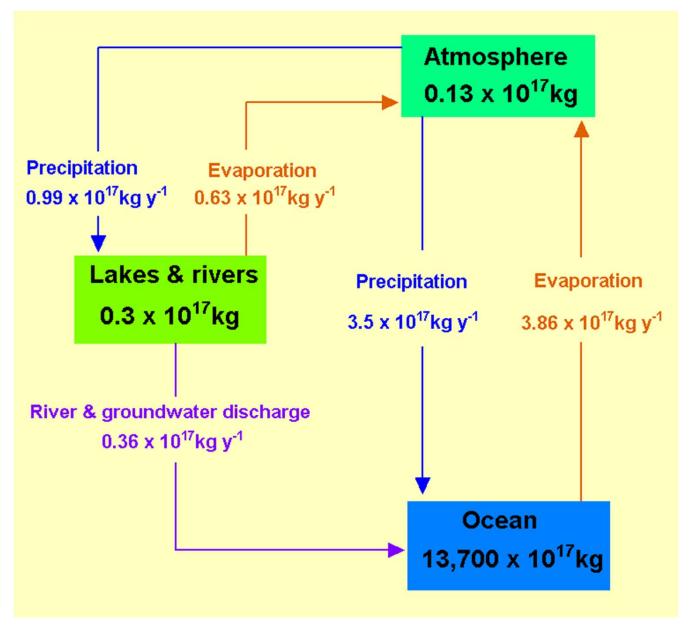


Figure 1-11. Simplified box model of the hydrologic cycle. Modified from Drever (1997).

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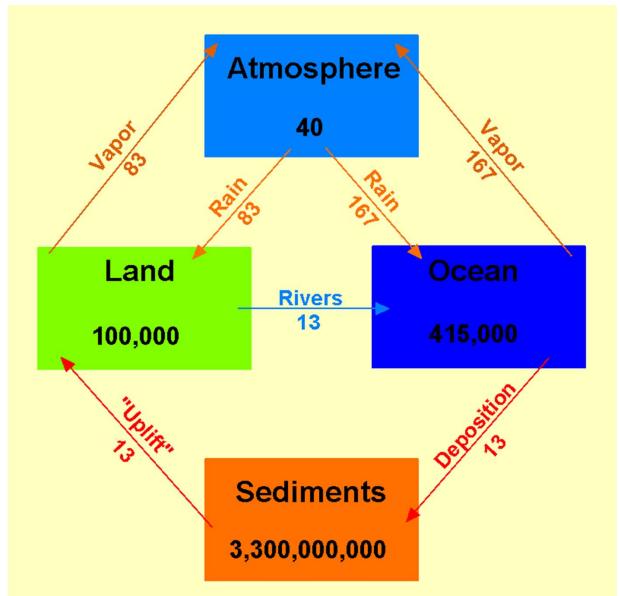


Figure 1-12. Prehuman cycle for mercury. Reservoir masses in units of 10^8 g. Fluxes in units of 10^8 g y⁻¹. From Garrels et al. (1975).

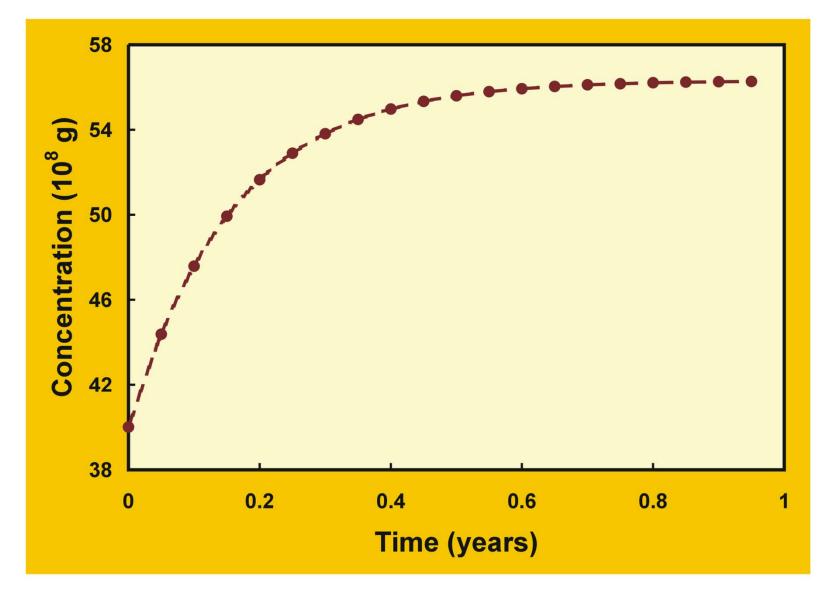


Figure 1-13. Variation in mercury content of the atmosphere as a function of time.

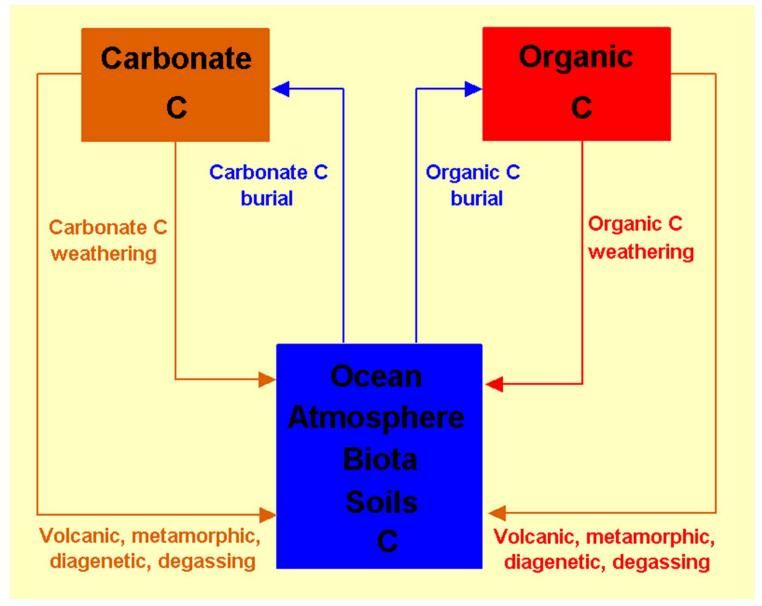


Figure 1-14b. The long term carbon cycle. After Berner (1999).

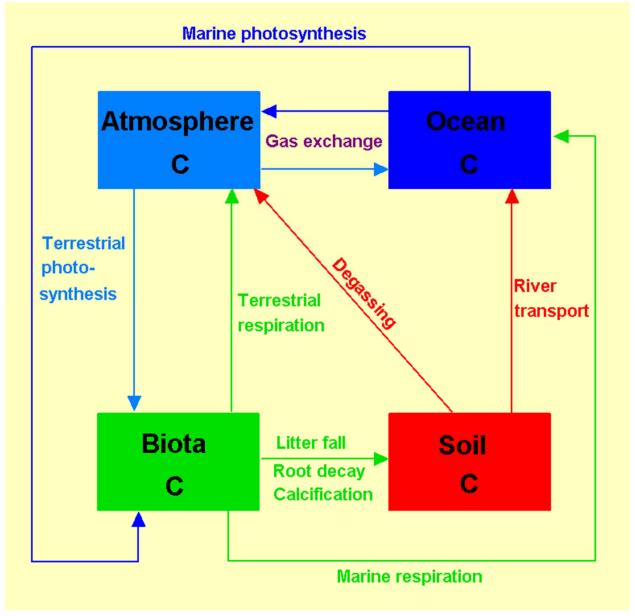


Figure 1-14a. The short-term carbon cycle, excluding anthropogenic inputs. After Berner (1999).